COPPER AVAILABILITY AS RELATED TO SOIL COPPER FRACTIONS IN OXISOLS UNDER LIMING

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ABSTRACT: The knowledge of the chemical forms of copper in soils and the relationships of these forms with soil copper availability are important for predicting the copper behavior in the soil-plant system. The present work studies the influence of liming on the available contents of copper as well as on the forms of copper fractions in six types of Oxisols. Soil samples, with and without liming, received copper at rates of 0.0, 20.0 and 40.0 mg dm³ and remained incubated for 30 days. Then, available copper was extracted with Mehlich-1, Mehlich-3, DTPA and EDTA solutions, and analyzed by atomic absorption spectrophotometry. Additionally, soil samples were extracted in a sequential procedure to determine Cu in fractions of soil, as follows: exchangeable-Cu fraction, organic matter-Cu fraction, Mn oxide-Cu fraction, amorphous Fe oxide-Cu fraction, crystalline Fe oxide-Cu fraction, residual-Cu fraction, and the total Cu content in the soil. Soil samples to which Cu was added presented higher Cu retention in the organic matter fraction with a small percentage retained in the exchangeable-Cu fraction. Liming resulted in a decrease of Cu in the exchangeable and organic matter fractions and an increase in the Fe and Mn oxide fractions and in the residual fraction. Without liming, the organic matter fraction presented the highest contribution to Cu content found in the soil extracts obtained with all extractors, except EDTA. For treatments with liming, Cu contents in the organic matter fraction were better correlated to Cu contents in extracts obtained with DTPA and Mehlich-3. Key words: fractioning, micronutrients, soil analysis

CALAGEM E OS TEORES DISPONÍVEIS E AS FRAÇÕES DE COBRE EM LATOSSOLOS

RESUMO: O conhecimento das formas químicas em que se encontra o cobre em solos, e suas relações com os teores disponíveis, são importantes para a previsão do seu comportamento no sistema solo-planta. Nesse sentido, este trabalho estuda a influência da calagem sobre os teores disponíveis e sobre o fracionamento de Cu em amostras de seis Latossolos. Essas amostras, submetidas ou não a calagem, receberam o elemento nas doses de 0,0, 20,0 e 40,0 mg dm⁻³ e foram incubadas por 30 dias. Terminada a incubação, procedeu-se a extração de Cu utilizando-se os extratores Mehlich-1, Mehlich-3, DTPA e EDTA. As amostras foram submetidas a um fracionamento que separou o Cu nas frações trocável, matéria orgânica, óxido de manganês, óxido de ferro amorfo, óxido de ferro cristalino e residual, além dos teores totais. Com sua aplicação ao solo o Cu foi retido, principalmente, na fração matéria orgânica, com pequena porcentagem retida na fração trocável. A calagem provocou redução nas frações trocável e matéria orgânica e aumento nas frações óxido de ferro, óxido de manganês e residual. Para os solos sem calagem, a fração matéria orgânica foi a que mais contribuiu para os teores determinados pelos extratores, com exceção do EDTA. Nos solos com calagem, o Cu ligado à matéria orgânica foi melhor correlacionado com os teores obtidos pelo DTPA e Mehlich-3. Palavras-chave: fracionamento, micronutrientes, análise de solo

INTRODUCTION

Initial efforts to establish methods for analysis of copper in soils focused the utilization of aqueous, diluted solutions of acids and bases, and diluted salts, which would simulate the root extraction power. These solutions were modified through the addition of anions and cations to displace exchangeable micronutrients, and of chelating agents which increase micronutrient solubility or compete with the organic matter for micronutrient complexation (Sims & Johnson, 1991).

have presented variable degree of success to make micronutrient deficiency or toxicity diagnoses for plants. The knowledge of the chemical forms in which micronutrients are extracted from the soil, for each extraction method, may help to understand the efficiency organic ohnson, To select a chemical solution which is able to

To select a chemical solution which is able to remove the nutrient quantity which represents their

Simple chemical extractions have been extensively used for the determination of micronutrient

availability in soils. However, as the amount of extracted

micronutrients varies with soil type, these extractions

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promptly and potentially available forms, the understanding of the various forms of the elements in the soils is required (Sims & Johnson, 1991). Therefore, methods of extraction which fractionate the soil and indicate the nutrient form more available to plants, as well as, the soil characteristics and management conditions affecting nutrient availability are very important for the prediction of nutrient behavior in the soil-plant system. A soil fractionation study which allows the indication of the nutrient form that better correlates to the amount of available nutrient determined by an extraction method is also important to predict nutrient behavior in the soil-plant system.

Soil pH is, most of the time, the main factor controlling Cu availability for plants, once increases of soil pH result in decreases of Cu availability. The negative effect of liming on Cu availability occurs, mainly, due to the increase in the soil cation exchange capacity which depends on the presence of pH dependent charges in the soil (Alloway, 1990). As the pH increases, the number of negative pH dependent charges also increases, the negative charge density at the surface of the colloid increases and the Cu availability for the plant decreases. Therefore, it is important to know how the pH affects the distribution of Cu in the different fractions of the soil (Shuman, 1985). The present work was carried out to study the influence of liming on the distribution of copper among the soil fractions. Additionally, the effect of liming was evaluated on the recovering rate of copper extracting the element from six Oxisols by chemical extractors.

MATERIAL AND METHODS

Samples from six soil types (0-20 cm layer), selected based on the basis of clay and organic matter contents, were collected at municipalities of the Minas Gerais State, Brazil (Table 1). They were: Rhodic Haplustox – LV1 (Sete Lagoas), Rhodic Haplustox – LV2 (São Sebastião do Paraíso), Rhodic Haplustox – LV3 (São Sebastião do Paraíso), Typic Haplustox – LV41 (João Pinheiro), Typic Haplustox – LVA2 (Viçosa) and Humic Haplustox – LVA3 (Viçosa).

Soil samples were homogenized and passed through a 2.0 mm sieve; sub-samples of 600 cm³ received the treatments. Half of the sub-samples were limed to adjust the pH to the 6.5 to 7.0 range, with rates of CaCO₃ and MgCO₃ based on an incubation study. After liming, soil samples were kept at field capacity for 15 days. Subsequently Cu rates were added (0.0, 20.0 and

Table 1 - Chemical and physical characterization of the soils (Rhodic Haplustox - LV1, Rhodic Haplustox - LV2, Rhodic Haplustox - LV3, Typic Haplustox - LVA1, Typic Haplustox - LVA2, Humic Haplustox - LVA3).

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|---|-------|-----------------|-----------------|------------------|------------------|------------------|
| Characteristic | LV 1 | LV ₂ | LV ₃ | LVA ₁ | LVA ₂ | LVA ₃ |
| рН Н,О (1:2,5) | 4.6 | 4.4 | 4.7 | 4.0 | 4.1 | 4.2 |
| P (mg dm ⁻³) ¹ | 4.0 | 0.0 | 0.0 | 1.3 | 0.0 | 1.9 |
| K (mg dm ⁻³) ¹ | 20 | 59 | 16 | 21 | 15 | 15 |
| Ca ²⁺ (cmol_dm ⁻³) ² | 1.24 | 0.55 | 0.72 | 0.08 | 0.02 | 0.46 |
| Mg ²⁺ (cmol _c dm ⁻³) ² | 0.20 | 0.11 | 0.23 | 0.04 | 0.02 | 0.13 |
| Al ³⁺ (cmol dm ⁻³) ² | 0.50 | 0.10 | 0.20 | 0.90 | 1.20 | 1.90 |
| H + AI (cmol dm ⁻³) ³ | 8.3 | 11.9 | 6.6 | 6.6 | 8.9 | 7.6 |
| Zn (mg dm ⁻³) ¹ | 1.21 | 0.46 | 0.37 | 0.40 | 0.18 | 1.14 |
| Fe (mg dm³) ¹ | 46.3 | 53.7 | 47.4 | 74.1 | 108.7 | 80.1 |
| Cu (mg dm³) ¹ | 0.77 | 6.57 | 8.73 | 0.40 | 0.12 | 0.22 |
| Mn (mg dm ⁻³) ¹ | 29.0 | 39.0 | 34.0 | 6.3 | 2.3 | 6.1 |
| CECefetiva (cmol_dm ⁻³) | 2.45 | 0.91 | 1.19 | 1.07 | 1.28 | 2.53 |
| CECtotal (cmol _c dm ⁻³) | 10.25 | 12.71 | 7.59 | 6.77 | 8.98 | 8.23 |
| V (%) | 14.5 | 6.4 | 13.0 | 2.5 | 0.87 | 7.7 |
| m (%) | 20.4 | 11.0 | 16.8 | 84.1 | 93.81 | 75.1 |
| O.C. (g kg ⁻¹) ⁵ | 28.7 | 13.92 | 17.0 | 8.7 | 18.6 | 112.5 |
| P-rem (g kg ⁻¹) ⁶ | 7.48 | 3.61 | 6.54 | 28.65 | 2.76 | 1.34 |
| SiO ₂ (g kg ⁻¹) ⁷ | 183 | 48 | 74 | 128 | 198 | 102 |
| Al ₂ O ₃ (g kg ⁻¹) ⁷ | 318.0 | 341.1 | 321.0 | 168.7 | 288.5 | 190.5 |
| Fe ₂ O ₃ (g kg ⁻¹) ⁷ | 124.2 | 338.4 | 304.1 | 46.0 | 126.6 | 53.7 |
| TiO ₂ (g kg ⁻¹) ⁷ | 7.1 | 67.5 | 61.5 | 5.1 | 19.3 | 6.4 |
| MnO (g kg ¹) ⁷ | 0.3 | 1.5 | 1.5 | 0.1 | 0.1 | 0.1 |
| $P_{2}O_{5} (g kg^{-1})^{7}$ | 2.1 | 2.4 | 2.2 | 0.3 | 0.6 | 1.1 |
| Thick sand(g kg ⁻¹) ⁸ | 110 | 70 | 90 | 400 | 150 | 460 |
| Fine sand (g kg ⁻¹) ⁸ | 40 | 230 | 300 | 230 | 100 | 140 |
| Silt (g kg ⁻¹) ⁸ | 140 | 240 | 190 | 50 | 70 | 160 |
| Clay (g kg ¹) ⁸ | 710 | 460 | 420 | 320 | 680 | 240 |

¹Mehlich-1 (De Filippo & Ribeiro, 1997). ²KCl 1 mol L⁻¹ (De Filippo & Ribeiro, 1997). ³CaOAc 0.5 mol L⁻¹ (De Filippo & Ribeiro, 1997). ⁴Tri-acid attack (Ure, 1990). ⁸Walkley-Black (De Filippo & Ribeiro, 1997). ⁶Remaining-P (Alvarez V. et al., 2000). ⁷Sulfuric digestion (Vettori, 1969). ⁸EMBRAPA (1997).

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40.0 mg dm⁻³) as $CuCl_2.2H_2O$ and samples were maintained incubated, at field capacity, for others 30 days.

After 30 days incubation soil samples were again passed through a 2 mm sieve and extracted with Mehlich-1 (De Filippo & Ribeiro, 1997), Mehlich-3 (Mehlich, 1984), DTPA (Lindsay & Norvell, 1978) and EDTA (Lantmann & Meurer, 1982). Copper in the extracts was determined by atomic absorption spectrophotometry.

Additionally, extraction was performed sequentially, based on Shuman (1985), except the amorphous Fe oxide fraction which was extracted based on Chao & Zhou (1983). This sequential fractionation separated Cu in the following fractions: exchangeable, bounded to organic matter, to manganese oxide, amorphous iron oxide and crystalline iron oxide. The total contents of micronutrients was determined after acid digestion (Ure, 1990). The sequential fractionation method used consisted, in detail, of:

Exchangeable Fraction (Ex) – In 50 mL centrifuge tube, a mixture of 5.0 g of soil and 20.0 mL of 1 mol L⁻¹ Mg(NO₃)₂ was shacken for two hours; the mixture was centrifuged for 10 min at 3000 rpm and the supernatant filtered and separated; 20.0 mL of distilled water were added, another shacking procedure was performed during 3 min, a new centrifugation was performed and the supernatant filtered and separated again. The two supernatants were combined and stored for analysis. The sequential extraction continued in the remaining of the soil sample.

Organic Matter Fraction (OM) – To the Ex free soil sample 10 mL of NaClO 5-6% at pH 8.5 were added; the system was heated to 100°C during 30 min., shacking from time to time; the sample was, thereafter, centrifuged and the supernatant filtered. This procedure was repeated twice and the three separated supernatants were combined. To the sample in the centrifuge tube, 10 mL of distilled water were added and after shacking for 3 min the mixture was centrifuged, filtered, and the separated supernatant was added to the three combined extracts from the NaCIO extraction and stored for analysis. The sequential extraction continued in the remaining soil sample.

Manganese Oxide Fraction (MnOx) – To the remaining soil sample 30.0 mL of NH₂OH.HCl (hydroxylamine hydrochloride) 0.1 mol L⁻¹, at pH 2.0 were added, and the mixture was shacken for 30 min, centrifuged and filtered; the separated supernatant was stored for analysis. To continue the sequential extraction, 15 mL of distilled water were added to the centrifuge tube, shacken for 3 min and the the supernatant discharged.

Amorphous Iron Oxide Fraction (AFeOx) – To the OM free soil sample 30.0 mL of NH ₂OH.HCl (hydroxylamine hydrochloride) 0.1 mol L⁻¹ plus HCl 0.25 mol L⁻¹, at pH 3.0 were added, and the mixture was shacken for 30 min, centrifuged and filtered; the separated supernatant was stored for analysis. To

continue the sequential extraction, 15 mL of distilled water were added to the centrifuge tube, shacken for 3 min and the supernatant discharged.

Crystalline Iron Oxide Fraction (CFeOx) – To the AFeOx free soil sample 30.0 mL of $(NH_4)_2C_2O_4$ (ammonium oxalate) 0.2 mol L⁻¹ + H_2C_2O_4 (oxalic acid) 0.2 mol L⁻¹ + ascorbic acid 0.01 mol L¹, at pH 3.0 were added, the mixture was heated to 100°C for 30 min, shacking from time to time; thereafter centrifuged and filtered; the separated supernatant was stored for analysis.

Residual Fraction (R) – Copper contents in the residual fraction were calculated by the difference between the determined total content and the summation of the contents found in the former fractions.

Total Content (T) – Total content of Cu was determined after acid digestion of soil samples (Ure, 1990), without sequential fractionation.

The experiment was set in an entirely randomized design with six soil types, three copper rates, with and without liming, and three replications. Correlation coefficients were made for the relationship between the Cu contents obtained by the extractors and the Cu contents found in the soil fractions. Additionally, correlation coefficients were calculated for the relationship between soil characteristics and the slopes of the Recovered Cu/Added Cu ratio, for each extractor. The coefficients were tested at the 5 and 1%.

RESULTS AND DISCUSSION

The Cu contents originally found in the soils (Table 2) were slightly above the Cu concentrations usually expected in soils which are in the 2 to 100 mg kg⁻¹ range (Baker, 1990). These contents are related to the parent material and are found in higher contents in soils originated from mafic rocks (LV1, LV2 e LV3) and in lower contents in soils from granite or gneiss (LVA1, LVA2, LVA3).

For soil samples that received no Cu and no liming, Cu concentrations were present predominantly (65 to 88%) in the Residual fraction (R) (Table 2). In the remainder fractions the distribution of retained Cu was 7.5% in the Crystalline Fe Oxides fraction (CFeOx), 7.3% in the Organic Matter fraction (OM), and 3.5% in the Amorphous Fe Oxides fraction (AFeOx) (Table 2). Similar results for Cu distribution in soils were found by McLaren & Crawford (1973a) and Sims (1986).

The Cu concentrations of the Exchangeable fraction in the soils without liming and with Cu addition (Table 2) at soil pH below 5.0 (Table 1) were low. These results indicate a high Cu adsorption capacity of the soils and are in agreement with Msaky & Calvet (1990) that observed complete adsorption of the Cu added to soil, even at low pH. On the other hand, the elevation of soil pH above 5.0 resulted in decrease of Cu concentration in solution (Sanders, 1982), which indicates the influence of higher pH values on the Cu sorption to the soil. The metals with higher electronegativities form stronger covalent

bounds with oxygen atoms from minerals and, therefore, are preferentially adsorbed to them. The electronegativity values for Cu (1.90), Zn (1.65) and Mn (1.55) explain the higher selectivity of the soil minerals for Cu and the high hysteresis for the adsorption of this element to soil, as compared to the other (McBride, 1994).

For treatments without liming the Cu distribution corresponded to 17% in the OM fraction, 11% in the CFeOx fraction, 8% in the AFeOx fraction, 3% in the MnOx fraction, and only 1.4% in the Ex fraction (Table 2). The lowest value for Cu in the exchangeable form indicates a low affinity of soil cation exchange sites for Cu ions, as

Table 2 - Copper concentrations in the Exchangeable (Ex), Organic Matter (OM), Manganese Oxides (MnOx), Amorphous Fe Oxides (AFeOx), Crystalline Fe Oxides (CFeOx), Residual (R) fractions and Total content (T), and the relative proportion of Cu in each fraction as related to the total Cu content, for three Cu rates, with and without liming.

| 0 | Rate | E | х | 0 | Μ | Mn | Эх | AFe | Ox | CFe | Ox | R | | Т |
|------------------|---------------------|---------------------|-----|---------|----|---------|----------|---------|----|---------|----|---------|-----|--------------------|
| 5011 | mg dm ⁻³ | mg kg ⁻¹ | % | mg kg⁻¹ | % | mg kg⁻¹ | % | mg kg⁻¹ | % | mg kg⁻¹ | % | mg kg-´ | ' % | Mgkg ⁻¹ |
| | | | | | | Wit | hout L | iming | | | | | | |
| LV ₁ | 0.0 | 0.0 | 0 | 1.4 | 3 | 0.0 | 0 | 1.4 | 3 | 2.3 | 5 | 38.2 | 88 | 43.3 |
| | 20.0 | 0.0 | 0 | 10.5 | 15 | 0.7 | 1 | 7.7 | 11 | 5.9 | 9 | 44.7 | 64 | 69.5 |
| | 40.0 | 0.4 | 0 | 21.0 | 23 | 1.9 | 2 | 13.9 | 15 | 11.3 | 12 | 43.4 | 47 | 91.9 |
| LV_2 | 0.0 | 0.0 | 0 | 4.1 | 4 | 2.5 | 2 | 4.1 | 4 | 6.4 | 6 | 85.7 | 83 | 102.8 |
| | 20.0 | 0.0 | 0 | 10.2 | 9 | 6.4 | 6 | 7.7 | 7 | 8.5 | 7 | 81.9 | 71 | 114.8 |
| | 40.0 | 1.4 | 1 | 20.1 | 14 | 1.8 | 8 | 12.3 | 8 | 10.3 | 7 | 91.8 | 62 | 147.8 |
| LV ₃ | 0.0 | 0.2 | 0 | 7.9 | 6 | 2.2 | 2 | 5.1 | 4 | 7.9 | 6 | 101.9 | 81 | 125.4 |
| - | 20.0 | 1.3 | 1 | 13.6 | 9 | 3.8 | 2 | 7.9 | 5 | 9.8 | 6 | 116.0 | 6 | 152.3 |
| | 40.0 | 1.8 | 1 | 21.7 | 12 | 6.4 | 4 | 11.1 | 6 | 11.3 | 6 | 126.0 | 71 | 178.3 |
| LVA_1 | 0.0 | 0.0 | 0 | 1.1 | 7 | 0.0 | 0 | 0.3 | 2 | 0.9 | 6 | 14.1 | 86 | 16.5 |
| | 20.0 | 3.3 | 9 | 12.6 | 35 | 3.1 | 9 | 2.1 | 6 | 3.3 | 9 | 12.0 | 33 | 36.4 |
| | 40.0 | 3.8 | 7 | 28.4 | 56 | 3.7 | 7 | 3.7 | 7 | 3.6 | 7 | 7.6 | 15 | 50.7 |
| LVA ₂ | 0.0 | 0.1 | 1 | 2.7 | 12 | 0.0 | 0 | 1.1 | 5 | 1.5 | 7 | 16.3 | 75 | 21.7 |
| | 20.0 | 0.9 | 2 | 6.2 | 15 | 0.8 | 2 | 7.2 | 18 | 6.8 | 17 | 19.2 | 47 | 41.1 |
| | 40.0 | 1.3 | 2 | 13.6 | 24 | 1.1 | 2 | 12.5 | 22 | 10.1 | 17 | 19.1 | 33 | 57.7 |
| LVA ₃ | 0.0 | 0.0 | 0 | 1.4 | 12 | 0.5 | 4 | 0.4 | 3 | 1.8 | 15 | 7.6 | 65 | 11.7 |
| | 20.0 | 0.1 | 0 | 13.3 | 26 | 1.5 | 3 | 4.7 | 9 | 14.8 | 29 | 17.5 | 34 | 52.0 |
| | 40.0 | 0.4 | 0 | 25.9 | 27 | 3.0 | 3 | 10.1 | 11 | 27.4 | 29 | 27.6 | 29 | 94.5 |
| Mean | | 0.8 | 1.4 | 12.0 | 17 | 2.7 | 3 | 6.3 | 8 | 8.0 | 11 | 48.4 | 59 | 78.2 |
| C.V.(% | б) | 10.0 | | 9.7 | | 5.6 | | 7.9 | | 8.6 | | 7.5 | | 4.7 |
| | | | | | | W | /ith Lin | ning | | | | | | |
| LV ₁ | 0.0 | 0.0 | 0 | 1.7 | 4 | 0.0 | 0 | 1.8 | 4 | 1.7 | 4 | 38.1 | 88 | 43.3 |
| | 20.0 | 0.0 | 0 | 10.2 | 15 | 0.9 | 1 | 7.5 | 11 | 5.1 | 7 | 45.8 | 66 | 69.5 |
| | 40.0 | 0.2 | 0 | 17.5 | 19 | 2.9 | 3 | 15.4 | 17 | 9.3 | 10 | 46.6 | 51 | 91.9 |
| LV , | 0.0 | 0.0 | 0 | 1.4 | 1 | 4.2 | 4 | 4.5 | 4 | 3.2 | 3 | 89.5 | 87 | 102.8 |
| | 20.0 | 0.0 | 0 | 3.9 | 3 | 10.0 | 9 | 8.2 | 7 | 4.8 | 4 | 87.9 | 77 | 114.8 |
| | 40.0 | 0.0 | 0 | 4.8 | 3 | 17.6 | 12 | 9.8 | 7 | 6.3 | 4 | 109.3 | 74 | 147.8 |
| LV ₃ | 0.0 | 0.0 | 0 | 3.5 | 3 | 5.4 | 4 | 3.9 | 3 | 8.2 | 7 | 104.4 | 83 | 125.4 |
| | 20.0 | 0.2 | 0 | 8.0 | 5 | 10.6 | 7 | 7.4 | 5 | 11.1 | 7 | 115.0 | 76 | 152.3 |
| | 40.0 | 0.4 | 0 | 15 | 8 | 16.9 | 9 | 9.3 | 5 | 14.7 | 8 | 122.0 | 68 | 178.3 |
| LVA ₁ | 0.0 | 0.0 | 0 | 0.4 | 2 | 0.7 | 4 | 0.4 | 2 | 0.2 | 1 | 14.8 | 90 | 16.5 |
| | 20.0 | 0.4 | 1 | 7.0 | 19 | 7.1 | 20 | 5.1 | 14 | 0.9 | 2 | 15.9 | 44 | 36.4 |
| | 40.0 | 0.9 | 2 | 13.9 | 27 | 12.1 | 24 | 10.6 | 21 | 1.6 | 3 | 11.6 | 23 | 50.7 |
| LVA ₂ | 0.0 | 0.0 | 0 | 0.1 | 0 | 0.1 | 0 | 1.4 | 6 | 0.8 | 4 | 19.3 | 89 | 21.7 |
| | 20.0 | 0.4 | 1 | 2.6 | 6 | 1.4 | 3 | 6.5 | 16 | 8.3 | 20 | 21.9 | 53 | 41.1 |
| | 40.0 | 0.5 | 1 | 5.4 | 9 | 3.4 | 6 | 11.6 | 20 | 13.5 | 23 | 23.3 | 40 | 57.7 |
| LVA ₃ | 0.0 | 0.0 | 0 | 1.1 | 9 | 0.2 | 2 | 0.2 | 2 | 2.3 | 20 | 7.9 | 68 | 11.7 |
| | 20.0 | 0.3 | 1 | 9.0 | 17 | 1.3 | 3 | 1.9 | 4 | 16.4 | 32 | 23.1 | 44 | 52.0 |
| | 40.0 | 0.3 | 0 | 19.4 | 21 | 2.3 | 2 | 5.1 | 5 | 46.2 | 49 | 21.2 | 22 | 94.5 |
| Mean | | 0.2 | 0 | 6.9 | 10 | 5.4 | 6 | 6.1 | 9 | 8.6 | 12 | 51 | 63 | 78.2 |
| C.V.(% | <u>()</u> | 17.5 | | 5.5 | | 12.0 | | 6.7 | | 6.9 | | 7.0 | | 4.7 |

LV₁ - Rhodic Haplustox; LV₂ - Rhodic Haplustox; LV₃ - Rhodic Haplustox; LVA₁ - Typic Haplustox; LVA₂ - Typic Haplustox; LVA₃ - Humic Haplustox

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also observed by Atanossova & Okazaki (1997). The importance of organic matter as the main reservoir of potentially available Cu for plants was reported (McLaren & Crawford, 1973a; Shuman, 1985; Sims, 1986). It seems that the OM fraction, in equilibrium with the AFeOx and CFeOx fractions, is responsible for supplying Cu to the soil solution under intensive cultivation, when the existing Cu is insufficient for maintaining adequate plant growth. The organic matter seems to be a source of specific sites for Cu complexation in the soil (McLaren & Crawford, 1973a, b), due to the affinity of the Cu ion to form inner sphere complexes with humic substances (McBride, 1994).

The increase in pH due to liming resulted in a reduction of Cu in the Exchangeable (Ex) fraction and, mainly, in the Organic Matter (OM) fraction, with Cu being transferred to Mn and Fe Oxide fractions, which retained 27% of the Total Cu (Table 2). Similar results were observed by Jarvis (1981) and Sims (1986) and were, probably, due to a specific adsorption process which would cause a decrease in Cu availability. These results for Cu content in the fractions suggest an equilibrium for Cu distribution in the Exchangeable, Organic Matter, and Oxide fractions, which would be modified with Cu being transferred to the more available fractions Ex and OM, as pH is reduced. The Mn Oxide fraction presented. proportionally, the higher increment in the Cu adsorption. Manganese oxides have particularly high selectivity for Cu which indicates that covalent bonds have important contribution for Cu adsorption (McBride, 1994). McLaren & Crawford (1973b) concluded that Mn oxides were the dominant mineral constituents in the specific adsorption of Cu by soils. For more weathered soils, however, the Fe oxides are found in higher concentrations and become, quantitatively, more important for Cu adsorption (Table 2).

An important factor to be considered concerning the Cu retention by soil fractions is the reaction time between the element and the soil. A more extended time of equilibrium with the soil would result in more stable forms of Cu in the presence of crystalline oxides since the reaction could continue slowly due to the diffusion of adsorbed ions into the interior of the particles (Barrow, 1993). In the present work this was observed for the soils which did not receive liming nor Cu addition (Table 2). These results demonstrate the importance of the time in the transformations of the added Cu to the soils and indicate that in soils fertilized with Cu the residual effect of the fertilizer decreases with its residing time in the soil, as observed by Levesque & Mathur (1986). This effect is incremented by liming.

In soils that did not receive lime, the extractors Mehlich-1, Mehlich-3 and DTPA yielded low contents of Cu from the humic LVA3 soil (Table 3) which indicates a high stability of the linking between Cu and Organic matter. On the other hand, in treatments with liming, at higher pH, the elevated contents of Cu extracted by DTPA and EDTA suggest that these extractors compete strongly with the organic matter for complexation of Cu, on these conditions. This happens because the stability of the

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complexes micronutrient-chelates is highly pH dependent (Inskeep & Baham, 1983; Stevenson & Fitch, 1989).

There was no correlation between the clay contents of the soils and the slopes of the Cu recovered/Cu added curves, since the extractions were performed in the soils without liming (Table 4). This is, probably, due to differences in organic matter content among the soils and to the higher affinity between Cu and organic compounds as compared to the mineral surfaces, as observed by Narwal & Singh (1998). The remaining-P (Alvarez V. et al., 2000) is a measurement that may be used as a parameter to evaluate the clay quantity and clay quality of the soil and has good correlation with the adsorption process. Therefore, the slopes of the Cu recovered/Cu added curves may be better correlated to the remaining-P than to the clay content which is only quantitative. In limed soils, the slopes of the Cu recovered/Cu added curves were negatively correlated to Cu content extracted by Mehlich-3, DTPA and EDTA, but not with those extracted by Mehlich-1. This indicates that the recovering rate of Cu by Mehlich-1 presented low sensibility in relation to soil clay content.

The susceptibility of the soil fraction to the extractor, evaluated by correlation, shows, in general, that the smaller correlations are obtained for the less stable fractions (Crystalline Fe oxides and Residual) for all extractors (Table 5). This indicates that the extractors remove, preferentially, the Cu located in the more promptly available sites. Micronutrients linked to Crystalline Fe oxides are found, mostly, occluded (Shuman, 1988) whereas those related to the Residual fraction are found associated to silicated minerals, resistant sulphydes, and refractory organic materials (Narwal & Singh, 1998). The Crystalline Fe oxides and Residual fractions are chemically and biologically stable.

The Organic Matter fraction presented better correlation with Cu content extracted with Mehlich-3 and DTPA, with and without liming (Table 5) which indicates that, in some way, Cu is being dissolved mainly from this fraction by these extractors. These results corroborate to the statement of McLaren & Crawford (1973a) that organic matter is the main reservoir of Cu availability for plants. In limed soils, however, Mehlich-3 was more sensible for the transference of Cu to the Amorphous Fe Oxides and the Manganese Oxide fractions as compared to the DTPA extractor. Mehlich-1 did not present selectivity for Cu extraction from any particular fraction and the correlation coefficients were similar among the fractions for the soils without liming; for limed soils, Mehlich-1 extracted Cu linked, mainly, to amorphous Fe oxides and Mn oxides, probably, due to its high acidity. Mehlich-1 has been criticized by the fact that it might extract non available forms of nutrients linked to specific adsorption sites and its high acidity extracts Cu linked to hydroxides and carbonates which are not available to plants, facts that would limit its utilization for acid soils.

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Table 3 - Copper extracted by Mehlich-1 (M-1), Mehlich-3 (M-3), DTPA and EDTA from samples of six soil types from Minas Gerais, Brazil which received three rates of Cu, with and without liming, and the respective slopes of the recovered Cu / added Cu curves (mg dm⁻³/ mg dm⁻³).

| Cail | Rate | M· | -1 | M· | -3 | DT | | EDTA | |
|------------------|------|---------|-------|---------|--------------|---------------------|-------|---------|-------|
| 5011 | | mg dm⁻³ | Slope | mg dm-3 | Slope | mg dm ⁻³ | Slope | mg dm⁻³ | Slope |
| | | | | W | ithout Limii | ng | | | |
| LV ₁ | 0 | 0.96 | | 1.76 | | 1.31 | | 1.42 | |
| - | 20 | 10.83 | | 10.16 | | 8.98 | | 11.92 | |
| | 40 | 22.02 | 0.53 | 21.79 | 0.50 | 18.55 | 0.43 | 23.28 | 0.55 |
| LV ₂ | 0 | 7.62 | | 4.44 | | 5.55 | | 7.59 | |
| | 20 | 21.28 | | 14.27 | | 13.59 | | 18.86 | |
| | 40 | 40.81 | 0.83 | 24.79 | 0.51 | 24.50 | 0.47 | 35.66 | 0.70 |
| LV ₃ | 0 | 10.19 | | 6.89 | | 7.87 | | 11.21 | |
| | 20 | 29.67 | | 17.59 | | 16.54 | | 25.47 | |
| | 40 | 45.07 | 0.87 | 24.78 | 0.45 | 25.21 | 0.43 | 44.54 | 0.83 |
| LVA, | 0 | 0.57 | | 0.66 | | 0.62 | | 0.69 | |
| | 20 | 20.34 | | 14.61 | | 9.83 | | 14.64 | |
| | 40 | 41.71 | 1.03 | 34.18 | 0.84 | 22.67 | 0.55 | 28.25 | 0.69 |
| LVA ₂ | 0 | 0.46 | | 0.60 | | 0.28 | | 0.62 | |
| 2 | 20 | 9.80 | | 6.08 | | 5.17 | | 13.04 | |
| | 40 | 22.94 | 0.56 | 14.43 | 0.35 | 12.90 | 0.31 | 20.18 | 0.49 |
| LVA ₃ | 0 | 0.38 | | 0.31 | | 0.52 | | 0.47 | |
| | 20 | 4.91 | | 8.25 | | 3.43 | | 14.59 | |
| | 40 | 12.58 | 0.30 | 15.97 | 0.39 | 11.35 | 0.27 | 23.79 | 0.58 |
| | | | | | | | | | |
| Mean | | 16.79 | 0.69 | 12.31 | 0.51 | 10.49 | 0.41 | 15.84 | 0.64 |
| C.V. (%) | | 9.47 | | 13.30 | | 10.66 | | 13.04 | |
| | | | | · ١ | With Liming | g | | | |
| LV ₁ | 0 | 0.20 | | 1.15 | | 0.87 | | 2.14 | |
| | 20 | 7.14 | | 9.39 | | 5.18 | | 16.74 | |
| | 40 | 17.39 | 0.43 | 21.65 | 0.51 | 14.28 | 0.33 | 31.09 | 0.72 |
| LV ₂ | 0 | 6.66 | | 5.38 | | 3.74 | | 10.20 | |
| | 20 | 22.19 | | 14.78 | | 7.58 | | 26.73 | |
| | 40 | 38.66 | 0.80 | 24.22 | 0.47 | 17.29 | 0.34 | 45.82 | 0.89 |
| LV ₃ | 0 | 9.46 | | 6.34 | | 5.64 | | 10.23 | |
| 5 | 20 | 27.30 | | 16.96 | | 14.11 | | 28.16 | |
| | 40 | 39.10 | 0.74 | 32.45 | 0.62 | 27.59 | 0.55 | 45.46 | 0.88 |
| LVA, | 0 | 1.49 | | 1.15 | | 0.29 | | 0.62 | |
| | 20 | 25.49 | | 14.58 | | 8.28 | | 22.09 | |
| | 40 | 37.33 | 0.90 | 29.70 | 0.71 | 16.49 | 0.40 | 47.98 | 1.18 |
| LVA | 0 | 0.23 | | 0.30 | | 0.35 | | 0.37 | |
| 2 | 20 | 7.47 | | 8.56 | | 3.00 | | 12.20 | |
| | 40 | 18.47 | 0.46 | 14.40 | 0.35 | 10.42 | 0.25 | 25.94 | 0.64 |
| LVA, | 0 | 0.18 | | 0.65 | | 0.75 | | 0.48 | |
| | 20 | 3.76 | | 10.29 | | 10.45 | | 15.26 | |
| | 40 | 10.81 | 0.27 | 23.52 | 0.57 | 24.85 | 0.60 | 39.80 | 0.98 |
| | | | | | | | | | |
| Mean | | 15.18 | 0.60 | 13.08 | 0.54 | 9.51 | 0.41 | 21.18 | 0.88 |
| C.V. (%) | | 14.24 | | 9.80 | | 14.02 | | 5.42 | |

LV₁ - Rhodic Haplustox; LV₂ - Rhodic Haplustox; LV₃ - Rhodic Haplustox; LVA₁ - Typic Haplustox; LVA₂ - Typic Haplustox; LVA₃ - Humic Haplustox

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Table 4 - Simple linear correlation coefficients between soil characteristics and slopes for the recovering curve obtained for Cu extracted by Mehlich-1 (M-1), Mehlich-3 (M-3), DTPA and EDTA, with and without liming.

| withouth | mmg. | | | |
|-------------------------|----------------------|----------------------|----------------------|----------------------|
| Charactristic | M-1 | M-3 | DTPA | EDTA |
| | | Withou | t Liming | |
| рН | 0.00 ^{NS} | - 0.33 ^{NS} | 0.08 ^{NS} | 0.80** |
| Organic Carbon | - 0.79** | - 0.41 ^{NS} | - 0.73** | 0.00 ^{NS} |
| Remanescent-P | 0.69** | 0.96** | 0.76** | - 0.56* |
| Clay | - 0.15 ^{NS} | - 0.34 ^{NS} | - 0.06 ^{NS} | - 0.08 ^{NS} |
| CEC | - 0.83** | - 0.34 ^{NS} | - 0.56 ^{NS} | - 0.14 ^{NS} |
| | | With | Liming | |
| рН | 0.60** | 0.27 ^{NS} | - 0.49* | 0.17 ^{NS} |
| Organic Carbon | - 0.76** | 0.02 ^{NS} | 0.64** | 0.12 ^{NS} |
| Remanescent-P | 0.64** | 0.68** | - 0.06 ^{NS} | 0.63** |
| Clay | - 0.20 ^{NS} | - 0.69** | - 0.79** | - 0.90** |
| CEC | - 0.03 ^{NS} | - 0.2* | - 0.48** | - 0.45 ^{NS} |
| ** *Ciencificant at the | 1 and 50/ " | بالمريكة ممرمم | NS Name alam | if a a st |

**, *Significant at the 1 and 5% respectively; [№] Non significant.

Table 5 - Simple linear correlation coefficients between Cu contents extracted by Mehlich-1 (M-1), Mehlich-3 (M-3), DTPA and EDTA and Cu in fractions Exchangeable (Ex), Organic Matter (OM), Manganese Oxides (MnOx), Amorphous Fe Oxides, Crystalline Fe Oxides, Residual (R), and Total (T) in six soil types, with liming (with) and without liming (without).

| Extractor | Liming | Ex | ОМ | MnOx | AFeOx | CFeOx | R | Т |
|-----------|---------|--------------------|--------|--------|--------|--------------------|--------|--------|
| M-1 | Without | 0.70** | 0.78** | 0.76** | 0.61** | 0.23 ^{NS} | 0.47** | 0.66** |
| | With | 0.35** | 0.46** | 0.93** | 0.69** | 0.06 ^{NS} | 0.47** | 0.62** |
| M-3 | Without | 0.71** | 0.91** | 0.65** | 0.61** | 0.35** | 0.28* | 0.54** |
| | With | 0.42** | 0.80** | 0.76** | 0.75** | 0.42** | 0.32** | 0.61** |
| DTPA | Without | 0.58** | 0.84** | 0.77** | 0.72** | 0.35* | 0.50** | 0.71** |
| | With | 0.33** | 0.83** | 0.66** | 0.58** | 0.65** | 0.36** | 0.66** |
| EDTA | Without | 0.16 ^{NS} | 0.53** | 0.67** | 0.75** | 0.54** | 0.59** | 0.75** |
| | With | 0.38** | 0.75** | 0.78** | 0.76** | 0.43** | 0.34** | 0.62** |

**, *Significant at the 1 and 5% respectively; NS Non significant.

extractor. Mehlich-1 did not present selectivity for Cu extraction from any particular fraction and the correlation coefficients were similar among the fractions for the soils without liming; for limed soils, Mehlich-1 extracted Cu linked, mainly, to amorphous Fe oxides and Mn oxides, probably, due to its high acidity. Mehlich-1 has been criticized by the fact that it might extract non available forms of nutrients linked to specific adsorption sites and its high acidity extracts Cu linked to hydroxides and carbonates which are not available to plants, facts that would limit its utilization for acid soils.

EDTA presented a similar behavior to Mehlich-1, including the possibility of extraction of non available forms of Cu due to its high concentration of compounds able to form complexes (Haynes & Swift, 1983) and the possibility of dissolution of Fe oxides (Borggaard, 1979).

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