

Soil hydrophobicity: comparative study of usual determination methods

Hidrofobicidade do solo: estudo comparativo dos métodos usuais de determinação

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ABSTRACT

Hydrophobic or water repellent soils slowly absorb water because of the low wett ability of the soil particles which are coated with hydrophobic organic substances. These pose significant effects on plant growth, water infiltration and retention, surface runoff and erosion. The objective of this study was to compare the performance of tension micro-infiltrometer (TMI) and the water drop penetration time (WDPT) methods in the determination of the hydrophobicity index of eighteen soils from southern Brazil. Soil samples were collected from the 0-5cm soil layer to determine particle size distribution, organic matter content, hydrophobicity index of soil aggregates and droplet penetration time of disaggregated and sieved soil samples. For the TMI method the soil samples were subjected to minor changes due to the use of macroaggregates to preserve the distribution of solid constituents in the soil. Due to the homogeneity of the soil samples the WDPT method gave smaller coefficients of variation unlike the TMI method where the soil structure is preserved. However, both methods had low coefficients of variation, and are thus effective for determining the soil hydrophobicity, especially when the log hydrophobicity index or log WDPT is >1.

Key words: water repellency, water drop penetration time, water infiltration.

RESUMO

Solos hidrofóbicos ou repelentes à água geralmente absorvem-na lentamente, devido à reduzida sortividade proveniente do recobrimento das partículas do solo por substâncias orgânicas hidrofóbicas. A hidrofobicidade do solo é uma propriedade com significativos efeitos no crescimento de plantas, infiltração e retenção de água, escoamento superficial e erosão. O objetivo deste estudo foi comparar o desempenho do método do microinfiltrômetro de tensão (TMI) e do método do tempo de penetração da gota (WDPT) na determinação do índice

de hidrofobicidade em dezoito classes de solos do sul do Brasil. Amostras de solo foram coletadas na camada de 0-5cm para a determinação da distribuição granulométrica, o conteúdo de matéria orgânica, índice de hidrofobicidade em agregados de solo e o tempo de penetração da gota em amostras de solo desagregadas e peneiradas. O método do microinfiltrômetro de tensão as amostras foram submetidas a menores alterações, devido ao uso de macroagregados, preservando a distribuição dos compostos no solo. Em função da homogeneização da amostra de solo, o método do WDPT pode apresentar menores coeficientes de variação, diferentemente do método do microinfiltrômetro em que a estrutura do solo é preservada. Contudo, ambos os métodos apresentaram baixos coeficientes de variação e, portanto, são efetivos na determinação da hidrofobicidade do solo, especialmente quando o log do índice de hidrofobicidade ou log WDPT é >1.

Palavras-chave: repelência à água, tempo de penetração da gota, infiltração de água.

INTRODUCTION

Soil hydrophobicity is an environmental problem of soil repellency to water that hampers soil wetting. This is a global phenomenon, which affects infiltration as well as soil water retention and plant growth. Hydrophobicity can be responsible for enhanced surface runoff, erosion and preferential flow (VOGELMANN et al., 2013b). Due to this high relevance, a great number of studies have been conducted on possible causes of water repellency and the results point to a variety of factors causing and influencing soil water repellency (SCHAUMANN et al., 2007; DOERR et al., 2007).

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Records of hydrophobic soils in Brazil are found mainly in areas cultivated with woody species and grasses (VOGELMANN et al., 2013a).

The occurrence of hydrophobicity is typically associated with the coating of soil particles by hydrophobic organic substances. These substances may be of various origins; most of them derived from local vegetation due to hydrophobic organic compounds in their chemical composition and are released either by deposition or decomposition of plant material (DOERR et al., 2007). Hydrophobicity is found to be related with other factors such as soil moisture (DOERR et al., 2007), pH (MATAIX-SOLERA et al., 2007), soil particle size (VOGELMANN et al. 2013b), incidence of fires (VOGELMANN et al., 2012), and presence of some fungi and bacteria species (SCHAUMANN et al., 2007).

The assessment of the hydrophobicity index has been a subject of research for decades. The method of water drop penetration time (WDPT), proposed by LETEY (1969), has been used to measure the persistence of water repellency in disaggregated soil samples (MATAIX-SOLERA et al., 2011). However, results have been inconsistent because of the varying degree of aggregate disruption. LEEDS-HARRISON et al. (1994) proposed the tension micro-infiltrometer method, which measures the soil hydrophobicity index in macro aggregates with the advantage of being a measure of hydrophobicity of soil with little or no aggregate disruption.

Both methods have been commonly used in the evaluation of hydrophobic soils, however no studies have compared the accuracy and inter-relationship of the results obtained by different methodologies. Therefore, the objective of this study was to compare the results of hydrophobicity index using the tension micro-infiltrometer method and water drop penetration time approach in eighteen different soil classes from southern Brazil.

MATERIALS AND METHODS

Description of study site and soil sample collection and analysis

A laboratory study was conducted during the summer of 2011 with soil samples from different soil classes from Rio Grande do Sul State, southern Brazil. Sampling sites had vegetation composed mainly of natural grassland with of vegetation on the occurrence of hydrophobicity. The main vegetation species at the sites sampled were *Andropogonlateralis*, *Axonopusaffinis*, *Paspalum* spp. and *Aristidalaavis*, which composed more than 80% of the vegetation in the sampled sites.

The soils of the area were classified according to the Brazilian System of Soil Classification (EMBRAPA, 2006) and the Soil Taxonomy developed by USDA (SOIL SURVEY STAFF, 2010) (Table 1). At each soil class, four undisturbed soil samples were collected at 0 – 5cm

Table 1 - Soil classification by the Soil Taxonomy (SOIL SURVEY STAFF, 2010) and Brazilian Soil Classification System (EMBRAPA, 2006) and textural class of the studied soils.

Soil taxonomy	Brazilian soil classification system	Textural class
Albaqualf1	PlanossoloHáplicoEutrófico	Sandy clay loam
Albaqualf2	PlanossoloHáplicoEutrófico	Sandy loam
Argiudoll	LuvissoloCrômicoPálico	Sandy clay loam
Dystrudept1	CambissoloHúmicoAlumínico	Clay
Dystrudept2	CambissoloHáplicoEutrófico	Silty clay
Endoaqualf1	GleissoloHáplicoDistrófico	Loam
Endoaqualf2	LuvissoloHáplicoÓrtico	Loam
Haplohumult	ArgissoloVermelho-AmareloAlumínico	Clay loam
Hapludalf	Argissolo Bruno-Acinzentado Alítico	Sandy loam
Hapludert	VertissoloEbânicoÓrtico	Silty clay
Hapludox1	Latossolo Bruno Alumínico	Clay
Hapludox2	LatossoloVermelhoDistrófico	Clay
Hapludox3	LatossoloVermelhoDistrófico	Clay
Hapludox4	LatossoloVermelhoDistroférico	Clay
Kandiudox	NitossoloVermelhoDistroférico	Clay
Paleudalf1	ArgissoloVermelhoDistrófico	Sandy loam
Paleudalf2	ArgissoloVermelhoDistrófico	Sandy loam
Paleudult	ArgissoloVermelhoDistrófico	Clay loam

soil layer using core sampler of known volume. In the same point and same soil layer were collected disturbed and undisturbed soil samples (soil blocks) to the soil hydrophobicity analysis.

Granulometric composition was determined by the pipette method (GEE & BAUDER, 1986), while soil organic carbon was determined using the NELSON & SOMMERS (1996) method. Soil organic matter was obtained from soil organic carbon content by multiplying the latter by 1,724, since it is assumed that organic carbon participates with 58% in the composition of soil organic matter (NELSON & SOMMERS, 1996).

The bulk density was obtained using the paraffin clod method described by BLAKE & HARTGE (1986). Soil particle density was determined following the method of GUBIANI et al. (2006). Total porosity (f) was calculated using the equation 1:

$$f = \left(1 - \frac{Bd}{Pd}\right) \times 100 \quad (1)$$

where: Bd is bulk density and Pd is particle density.

Tension micro-infiltrometer technique

The undisturbed soil samples were air-dried in the laboratory until the point of brittleness. Subsequently, the aggregates were separated manually by means of traction, with respect to the cleavage surface, resulting in spheroidal-shape aggregates of about 20mm in diameter. These aggregates were then subjected to soil sorptivity test using the tension micro-infiltrometer, with 15 repetitions for each soil.

This test consisted of an apparatus with a tube in which one end is connected to a reservoir with a liquid, either distilled water or ethanol (95% v/v), and the other end has a small sponge (4mm of diameter) in contact with the aggregate for 2 minutes, and the cumulative mass of water or ethanol which infiltrates the soil by capillary is recorded as the difference in initial and final weight of the reservoir of liquid using an analytical scale with an accuracy of 0.001g. Hydraulic pressure differences which could affect flow within the column of fluid in the reservoir and the infiltrometer were eliminated. Error due to evaporative loss during the short testing time was reduced by applying a thin layer of silicone oil to the surface of the water reservoir. The two different liquids were employed during the sorptivity tests due to their different density, viscosity, and contact angle with the soil.

Water sorptivity test was first conducted on undisturbed aggregates, and afterward these soil

aggregates were air-dried again and ethanol sorptivity test was performed on them.

The sorptivity (S) was obtained by equation 2:

$$S = \sqrt{\frac{Qf}{4br}} \quad (2)$$

where: Q corresponds to the measurement of liquid flow ($\text{mm}^3 \text{ s}^{-1}$); b is dependent on the parameter of the function of diffusion of water in the soil being taken as 0,55; f is the total porosity ($\text{mm}^3 \text{ mm}^{-3}$); r is infiltrometer tip radius (0.5mm). Q is obtained by the fluid infiltration rate in a circular area on the surface of the soil aggregate (LEEDS-HARRISON et al., 1994).

The hydrophobicity index or index of repellency (R) was evaluated by comparing the values of water sorptivity and ethanol sorptivity using equation (3) as suggested by TILLMAN et al. (1989).

$$S_{\text{water}} = \left[\frac{(\mu_e / \gamma_e)^{1/2}}{(\mu_a / \gamma_a)^{1/2}} \right] S_{\text{ethanol}} \quad (3)$$

where: μ_e is the viscosity of ethanol (95% v/v) at 20°C ($0.0012 \text{ N s m}^{-2}$); γ_e is the surface tension of ethanol (95% v/v) at 20°C (0.023 N m^{-1}); μ_a is the water viscosity at 20°C ($0.0010 \text{ N s m}^{-2}$); γ_a is the surface tension of water at 20°C (0.073 N m^{-1}).

Using these values, equation 3 is simplified into equation 4 (TILLMAN et al., 1989):

$$S_{\text{water}} = 1.95 S_{\text{ethanol}} \quad (4)$$

Thus, a further simplification of equation 4 by TILLMAN et al. (1989) yielded different values for different soils and were assigned the index, R as:

$$R = 1.95 \left[\frac{S_{\text{ethanol}}}{S_{\text{water}}} \right] \quad (5)$$

Water drop penetration time (WDPT) test

The disturbed soil samples were air-dried, crushed and sieved through a 2-mm sieve. The soil samples were later placed in Petri dishes (volume of 25 cm^3) for the test. The WDPT method consisted of applying three drops of water with a Pasteur pipette, and then measuring the time taken by the drops to penetrate the soil sample (KING, 1981). Each drop was released from a height of 10mm above the soil surface to minimize the impact with the surface, with 15 repetitions for each soil.

Statistical analysis

The normal distribution of the experimental data was tested using the Shapiro-Wilk analysis. The WDPT data did not follow a normal distribution

and were subjected to logarithmic transformation. All data were subjected to analysis of variance (ANOVA) at 5% level of probability. When the F test was significant, means of hydrophobicity index and WDPT were compared by Tukey.

The following statistical performance indices were used to evaluate the degree of precision and accuracy of the estimates: index of Willmott "d" (WILLMOTT et al., 1985) and correlation coefficient "r". The index "d" indicates the degree of deviation of the estimated values with the values observed (0-1), being 0 for no agreement and 1 for perfect agreement, whereas the correlation coefficient "r" determines the accuracy of each method and indicates the degree of dispersion of the points from the average.

RESULTS AND DISCUSSION

There was considerable variability in general soil properties and specifically in particle size distribution and soil organic matter content. The textural class of the soil series varied from sandy loam to clay (Table 2). The hydrophobicity was more severe in Albaqualf2, Dystrudept2 and Hapludert soils, which showed levels of organic matter of 18, 41 and 56g kg⁻¹, respectively, showing that the

occurrence of hydrophobicity may be associated with a wide range of soil organic matter content.

By contrast, Arguidoll and Haplohumult soils did not show hydrophobicity even with high levels of organic matter (41 and 46g kg⁻¹, respectively). This indicates that the amount of organic material is not the only factor responsible for soil hydrophobicity, thus confirming the results already found by VOGELMANN et al. (2012) and MATAIX-SOLERA et al. (2007) which shows that hydrophobicity in the soil is due to the nature of organic compounds and not only on the amount of organic matter.

The hydrophobicity indices obtained using the two methods showed little variation, thus we can assume that both were potentially suitable for the analysis of the hydrophobic character of soils. There was little data dispersion around the line of the equation fitted to the observed data (Figure 1) as evidenced by the high coefficient of determination ($R^2=0.87$). This shows a satisfactory estimation of the hydrophobicity of the soil with allow coefficient of variation (Table 3). These values and coefficient of variation are in agreement with the values observed by LEEDS-HARRISSON et al. (1994) and VOGELMANN et al. (2012, 2013a).

Table 2 - Soil particle size distribution, bulk and particle density, total porosity, and soil organic matter content in the of 0-5cm layer of the studied soils.

Soil	---Granulometric composition (g kg ⁻¹)*---			Bulk density (g cm ⁻³)	Particle density (g cm ⁻³)	Total porosity (cm ³ cm ⁻³)	Organic matter (g kg ⁻¹)
	Sand	Silt	Clay				
Albaqualf1	537	176	287	1,15	2,50	0,54	31
Albaqualf2	654	240	107	1,12	2,55	0,56	18
Arguidoll	447	248	304	1,23	2,50	0,51	41
Dystrudept1	295	268	438	1,08	2,54	0,57	30
Dystrudept2	113	441	446	1,21	2,55	0,53	41
Endoaqualf1	339	410	251	1,26	2,39	0,47	37
Endoaqualf2	396	474	130	1,17	2,40	0,51	25
Haplohumult	245	490	265	1,29	2,46	0,48	46
Hapludalf	665	196	139	1,14	2,51	0,55	24
Hapludert	164	435	401	1,25	2,55	0,51	56
Hapludox1	56	349	595	1,19	2,55	0,53	37
Hapludox2	393	191	416	1,15	2,52	0,54	27
Hapludox3	430	168	403	1,21	2,57	0,53	28
Hapludox4	42	338	620	1,10	2,55	0,57	35
Kandiudox	307	196	497	1,25	2,48	0,49	26
Paleudalf1	643	253	100	1,03	2,49	0,60	27
Paleudalf2	592	237	170	1,11	2,56	0,57	13
Paleudult	395	337	269	1,25	2,55	0,51	21

*sand = 2-0,05mm; silt = 0,002-0,05mm; clay = <0,002mm.

Table 3 - Values of water and ethanol sorptivity, hydrophobicity index and water drop penetration time (WDPT) in the soil layer of 0-5cm of the studied soils.

Soil	Water sorptivity		Ethanol sorptivity		Hydrophobicity index			WDPT		
	----- (mm s ^{-1/2}) -----							----- (s) -----		
Albaqualf1	0,043	±0,005*	0,088	±0,010	4,0	c	±0,4	46	d	±3,0
Albaqualf2	0,034	±0,005	0,152	±0,023	8,7	b	±0,6	180	c	±25
Argiudoll	0,015	±0,002	0,016	±0,002	2,1	ef	±0,3	2	f	±0,5
Dystrudept1	0,049	±0,005	0,053	±0,006	2,1	ef	±0,2	2	f	±0,4
Dystrudept2	0,013	±0,002	0,060	±0,008	9,0	b	±0,8	235	b	±20
Endoaqualf1	0,053	±0,007	0,087	±0,011	3,2	d	±0,4	23	e	±2,5
Endoaqualf2	0,077	±0,008	0,051	±0,006	1,3	f	±0,1	2	f	±0,2
Haplohumult	0,077	±0,012	0,067	±0,010	1,7	f	±0,2	2	f	±0,2
Hapludalf	0,054	±0,007	0,063	±0,008	2,3	ef	±0,3	3	f	±0,3
Hapludert	0,009	±0,001	0,050	±0,007	10,9	a	±0,9	800	a	±54
Hapludox1	0,088	±0,011	0,122	±0,016	2,7	de	±0,3	3	f	±0,2
Hapludox2	0,042	±0,005	0,034	±0,004	1,6	f	±0,2	2	f	±0,1
Hapludox3	0,055	±0,006	0,085	±0,009	3,0	d	±0,3	24	e	±5,0
Hapludox4	0,074	±0,011	0,053	±0,008	1,4	f	±0,2	2	f	±0,4
Kandiudox	0,052	±0,007	0,059	±0,008	2,2	ef	±0,3	1	f	±0,2
Paleudalf1	0,087	±0,011	0,067	±0,009	1,5	f	±0,2	2	f	±0,4
Paleudalf2	0,034	±0,005	0,035	±0,005	2,0	ef	±0,2	1	f	±0,1
Paleudult	0,068	±0,009	0,101	±0,013	2,9	d	±0,3	4	f	±0,2
C.v.					12,7			9,3		

Values with the same letters are not statistically difference at P<0.05 significant level. C.v.-coefficient of variation. *Standard deviation.

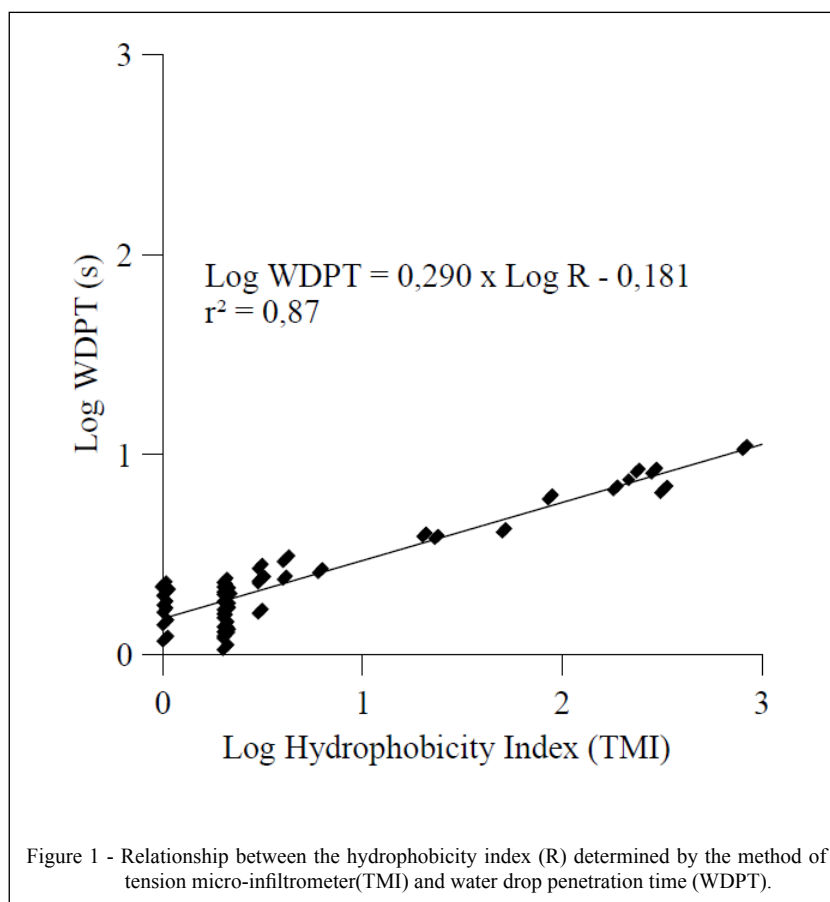
The lower coefficient of variation obtained by the method WDPT could be due to the fact that the soil samples were disturbed, that is, ground and sieved, so there may be a uniformity of hydrophobic compounds when present in the soil matrix. CERDÁ & DOERR (2008) described that the disintegration of the soil will expose internal areas of aggregates, thereby exposing areas that are not hydrophobic such as in the case of soils that were sieved and homogenized, which could result into reduction in hydrophobicity. VOGELMANN et al. (2012) reported that the compounds may accumulate on the outer surface of macroaggregates or in smaller clusters, which may experience a reduction in its concentration internally. Thus, the hydrophobicity is not uniformly distributed in the soil matrix. Hydrophobicity may be reduced by homogenizing the distribution of these compounds in the soil matrix.

In the tension of the micro-infiltrometer method, the samples experienced less disruption due to the use of macroaggregates. This treatment preserved the natural distribution of compounds in the soil matrix (LEEDS-HARRISON et al., 1994). Thus, the homogenization of hydrophobic compounds in the

soil matrix in the WDPT method may be responsible for the lower coefficient of variation obtained, as also observed by CERDÁ & DOERR (2008).

Moreover, it is noteworthy that the WDPT method does not consider all the original physical or hydraulic properties of soil which can restrict or favor the infiltration of water into the soil, like soil porosity and physical properties of liquids as is the case of viscosity and surface tension. However, these liquid properties are considered when using the tension micro-infiltrometer technique, as shown in equations 2 and 3. Failure of using such parameters may result in erroneous inference, for example, a soil with low porosity where water droplets do not infiltrate due to the small pore volume that restricts and decreases infiltration into the soil may be interpreted as a hydrophobic soil from the point of view of the WDPT method.

The index “d” of WILLMOTT et al. (1985), which reflects the accuracy of the method, indicated an “average” performance (0.57). As this index is a mathematical approximation in evaluating the correctness and dispersion of the values, it is evident that the results obtained by both methods show good



correlation between the results and a small scattering of data as shown in figure 1. The determination of soil hydrophobicity by both methods exhibited deviation when compared to each other, with high magnitude when either log hydrophobicity index or log WDPT is less than 1. This indicates a possible divergence in determining the degree of hydrophobicity in less hydrophobic soils. In contrast, for soils with high hydrophobicity degree, both methods gave greater similarity in the magnitudes of evaluation.

CONCLUSION

The methods of soil hydrophobicity determination through the use of tension micro-infiltrometer and the droplet penetration time (WDPT) gave low coefficients of variation and good correlation between the results, hence both were considered effective in evaluating soil hydrophobicity, especially when the log hydrophobicity index or log WDPT is >1 . As a result of the homogeneity of the soil samples, the WDPT method may provide lower

coefficients of variation because of the homogeneous distribution of hydrophobic compounds, in contrast to the tension micro-infiltrometer technique where the soil structure is preserved.

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