



Physical and chemical study of lattice kaolinites and their interaction with orthophosphate

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ABSTRACT

This study covers the evaluation of the structure and the capacity in adsorbing orthophosphate ions of kaolinites collected from sampling sites in Manaus (Brazil). The kaolinites were obtained using physical fractioning (sieving/siphoning) techniques and chemical treatment with HCl, H₂SO₄, H₂O₂ and KCl. The investigation of the kaolinite lattices involved the Hinckley and Plançon indexes determined from X-ray diffraction data, Fourier transformed infrared spectroscopy, scanning electron microscopy and Mössbauer spectroscopy. The absorption capacity of orthophosphate ions was calculated by Freundlich and Langergren isotherms. A transitional state was observed in the crystallographic structure from kaolinites because of the isomorphous substitution of Al³⁺ by Fe³⁺. This isomorphous substitution occurs accompanied by the optical pleochroism behavior, but it also reduces the mean particle sizes and increases the number of structural defects and magnetic properties of these kaolinites. Mössbauer spectroscopy showed that the substituting Fe³⁺ preferentially occupies octahedral sites. In the kaolinite lattices there are different octahedral sites of Al bounded by *cis*-OH-Fe³⁺ and *trans*-OH-Fe³⁺ octahedral sites. The kaolinite from the Kao1 sample has a higher number of *cis*-OH-Fe³⁺ from octahedral sites and is able to adsorb higher contents of orthophosphate ions than those from samples Kao2 and Kao3.

Key words: Freundlich, isomorphous substitution, Langergren, Manaus, Mössbauer.

INTRODUCTION

Kaolinite, a clay mineral formed by octahedral and tetrahedral sites (1:1), is the mineral that is most found in soil and sediment, mainly in humid tropical regions (Corrêa et al. 2008). This clay mineral is a typical 1:1-type phyllosilicate with the chemical composition Al₂Si₂O₅(OH)₄. Its surface comprises aluminol groups (≡Al-OH) situated at the edges and the hydroxyl-terminated planes of

the clay lamellae (Wei et al. 2014). On its basal surface there are permanent negative charges that are responsible for sorption of cations (Tombácz et al. 2004) and silanol (≡SiOH) linked to the aluminol groups through hydrogen atom (Rehim et al. 2010).

Anion exchange capacity (AEC) occurring on the edges of the aluminosilicate sand layer is pH-dependent. Kaolinite AEC increases with decreasing pH but never > 2.00 cmol kg⁻¹ (Dixon 1989). On its silicate layer, kaolinite notably absorbs phosphates. When linked on edge, the phosphates increases the

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reactivity of the kaolinite (Cui and Weng 2013). The pH value has great influence on the capacity adsorption of ions by kaolinite: i) pH <4.0-adsorbs anions, ii) pH ~4.0 is PCZ of the kaolinite it does not adsorb ions and iii) pH > 4.0 adsorbs cations (Khawmee et al. 2013). Specifically, the adsorption of phosphate occurs in pH 3.0-5.0 (Kamiyango et al. 2009) Malawi, has revealed high phosphate levels ranging from 0.63 to 5.50 mg/L. These phosphate levels would stimulate excessive growth of plants and toxic cyanobacteria in stagnant receiving water bodies hence posing a threat to aquatic life and water quality. Phosphate removal by kaolinite obtained from Linthipe, Malawi, was investigated as a function of pH, contact time, clay dosage, competing ions and initial phosphate concentration by means of jar tests. Phosphate uptake was pH dependent with adsorption mechanisms on kaolinite and iron oxide surfaces dominant generally below pH 7 and precipitation by calcium ions dominant above pH 7. Maximum phosphate removal occurred at high pH values of 11.22 (97.1% Obviously, the presence of substituents alters the particle size and crystallinity and consequently the ionic adsorption capacity of the kaolinite (Balan et al. 2000).

Since phosphorus is one of the essential elements required for plant growth kaolinite and its modified forms provide an important control over the availability of this nutrient to the soil solution. Phosphate adsorption onto minerals is an important process that significantly affects the mobility and bioavailability of phosphorus in natural environments. Furthermore, the kinetics and thermodynamics studies from the adsorption process provide important outcomes on available phosphorus. Previous researchers have reported numerous kinetic models used to describe sorption processes. These include first-order, and second-order reversible ones, and first-order and second-order, irreversible ones, pseudo-first-order and pseudo-second-order ones based on the solution concentration (Adebowale et al. 2008). The major

contribution of this study is to provide data on crystallinity and interaction with orthophosphate by kaolinites extracted from Oxisoil and bottom sediment from the Central Amazonia.

EXPERIMENTAL PART

SAMPLES COLLECTION

Samples of Oxisoil located in S 03°02'55.3" and W 60° 04' 21,4" (Kao1) and bottom sediments in S 03° 08' 45.9" and W 60° 01' 05,4" (Kao2) and S 03° 05' 53.9" and W 59° 58' 15" (Kao3). Soil samples were collected from 40 to 80 cm depths (C horizon), classified according to their physical and chemical characteristics in yellow latosols, and sediment samples were collected from the bottom of streams.

PREPARATION OF FINER KAOLINITE

To begin, all samples are dried at room temperature for a week and sieved down to a size < 0.053 mm to obtain finer particles from the mixed deionized water (~1 L) with the Kao1, Kao2 and Kao3 at < 0.053 mm under shaker agitation and kept for 24 hours, according recommended by Couceiro and Santana (1999). Then, the supernatant was siphoned and the residue dried at room temperature for 120 h and triturated in an agate mortar. Nearly 10.00 g of Kao1 with 100 mL of hydrochloric acid (1:1) were used in the treatment for the elimination of iron oxides such as hematite and goethite, a process which lasted for about 120 h. The separation of the residue occurred by centrifugation at 3,000 rpm (Damon/HN-SII) for 5 min and was washed thoroughly under manual shaking with demineralized water. About 10.00 g of Kao2 and Kao3, previously treated with hydrogen peroxide (30%), concentrated sulfuric acid and deionized water for removing organic matter receive the same separation treatment as Kao1. Aiming at maintaining the ionic equivalence among the samples, samples Kao1, Kao2 and Kao3 were kept

through a 12 h period of stirring with KCl 1.0 mol L⁻¹, dried at room temperature, crushed in an agate mortar and stored.

MINERALOGICAL COMPOSITION

To develop the study on the mineralogical composition of Kao1, Kao 2 and Kao3 the physical techniques XRD, FTIR and SEM were used. The XRD measurements carried out on a Shimadzu diffractometer LabX 6000 coupled to a source of Cu (K α) with $\lambda = 1.5418 \text{ \AA}$, angle ranging from 5° to 60° (2 θ min⁻¹), time constant at 5 s. The Hinckley and Plançon index was determined in the program Origin 8.0. The chemical analyses were carried out by X-ray fluorescence (spectrometer Philips PW2400). Operation of the X-ray spectrometer had been automated and goniometer angles, and kV and mA for samples.

Infrared measurements were carried out on KBr pellet dried at 100 °C at a ratio of 1:300 using a Fourier Transform Infrared (FTIR) spectrometer (model 640 FTIR, Varian). Conditions for obtaining spectra include a resolution of 4 cm⁻¹, scan number of 20 and a range of 4,000 to 400 cm⁻¹. The SEM electron microscope FEI QUANTA 250 in samples with deposition of a conductive tape, coated with a layer of gold and images recorded under accelerating voltage of 20 kV, a current of 6.0 x 10⁻¹¹ and 5x10⁻⁷ torr pressure. In addition, mapping of the elements was done by the Energy Dispersive System (EDS) EDAX model.

Iron was studied as kaolinite substituent with ⁵⁷Fe Mössbauer spectrometer at 298 and 80 K used in a conventional transmission assembly and a source of ⁵⁷Co/Rh. The correction of isomer shifts used α Fe. The data were numerically fitted with Lorentzian functions by the method of least squares, using the computer program NORMOS® 90, written by R. A. Brand, Laboratorium für Physik Argewandte, Universität Duisburg, D-47048, Duisburg-Germany.

ADSORPTION STUDY USING ISOTHERMS AND KINETICS

The adsorption isotherms for Kao1, and Kao2 Kao3 were carried out in flasks by tanking standard KH₂PO₄ solution with concentration at 3 mg L⁻¹ and pH ~ 4.0, adjusted with HNO₃ at 0.1 mol L⁻¹. According to Miranda-Trevino and Coles (2003) mainly as a result of the degree of weathering in the different compounds. Nevertheless, the kaolinite structure possesses great advantages in many processes due to its high chemical stability and low expansion coefficient. As a consequence of adsorption, the kaolinite structure and the soil solution pH will change. To analyze the adsorption behaviour of kaolinite, Pb, Zn and Cd were studied at three different concentrations (1, 2 and 3 mmol/l in pH 4.0 value the kaolinite superficial charge is altered to positive that is ideal for anions adsorption. Each isotherm comprised kaolinite mass from 0.050 to 0.500 g, ranging 0.05 g under agitation on shaker for 2 h and centrifugation at 3,000 rpm for 5 min. The supernatants were filtered on Millipore cellulose ester membrane with 45 mm in pore diameter and 47 mm (Santos et al. 2007). All the adsorption isotherms developed in triplicate.

The kinetic curves for Kao1, and Kao2 Kao3 were carried out by the flask tanking standard KH₂PO₄ solution with concentration at 3 mg L⁻¹, pH ~ 4 adjusted with HNO₃ to 0.1 mol L⁻¹. Each kinetic curve comprised kaolinite mass of 0.30 g under different times of agitation. The flasks were stirred in a shaker at intervals of 10 min and ranged from 0 to 120 min. After each interval, one flask was centrifuged at 3,000 rpm for 5 min. The supernatants were filtered on Millipore cellulose ester membrane with 45 mm in pore diameter and 47 mm (Santos et al. 2007). All the kinetic curves developed in triplicate.

The KH₂PO₄ contents of the supernatants obtained in the adsorption isotherms and kinetic curves were analyzed by ultraviolet-visible

spectrophotometry and the ammonium molybdate method at 885 nm.

The equations of Freundlich and Lagergren pseudo-second order adjusted to the results obtained by adsorption isotherms and kinetic curves (Dotto et al. 2011, Souza and Santana 2014):

1. Freundlich equation

$$q_e = k_f C_e^{1/n}$$

Where q_e is amount of phosphorus adsorbed, C_e is equilibrium concentration of phosphorus in solution, k_f and n are empirical constants.

2. Lagergren equation

$$\log_{10}(q_e - q) = \log_{10} q_e - k_f t / 2,303$$

Where q_e and q are phosphorus concentrations adsorbed in equilibrium in time t respectively, k_f is the adsorption velocity constant.

RESULTS AND DISCUSSION

Table I shows that the SiO_2 and Al_2O_3 contents determined in the Kao1, Kao2, and Kao3 are very different from theoretical values: 46.5 % - SiO_2 and 39.5 % - Al_2O_3 (Aranha et al. 2010). This difference is explained by the high level weathering Amazonian region condition. Additionally, Fe_2O_3 , TiO_2 , and in minor proportion, MgO , and CaO were identified. The literature shows that mineral impurities such as quartz, anatase, rutile, pyrite, limonite, feldspar, mica, montmorillonite, and various iron and titanium oxides in kaolinites are usually found. Si and Al, in the form of hydroxides, can apparently occur as coatings on the kaolinite layers. In fact, the contents of SiO_2 , Al_2O_3 and Fe_2O_3 correspond to kaolinite, quartz and gibbsite (XRD not shown). The presence of Fe and $\text{SiO}_2/\text{Al}_2\text{O}_3$ ratio indicate that Kao1, Kao2, and Kao3 can have isomorphic substitution of Al^{3+} by Fe^{3+} (Melo et al. 2002). The literature explained the TiO_2 presence as an aggregate commonly found in kaolinite. The TiO_2 content of kaolinite is apparently dependent

on the mineral source. For example, feldspar, such as kaolinite from Manaus, has a low TiO_2 content compared to a biotite schist or biotite granite source which shows a relatively high TiO_2 (Weaver and Pollard 1973).

Figure 1 clearly shows that d_{011} , $d_{01\bar{1}}$, $d_{0\bar{1}\bar{1}}$ values are distinguishable for the three kaolinites. The kaolinite XRD presents (hkl) reflections commonly affected by structural defects. Among such reflections, we can distinguish those wherein k is multiple of 3, i.e., $k = 3n$ ($n = 1, 2, 3, \dots$) and those wherein $k \neq 3n$. The existence within pilings of defects caused by displacements and/or rotations at the plane layers explains the multiple observed k . Therefore, the Hinckley and Liètard indices indicate that three kaolinites have important structural differences (Table II). Typically the modifications occur in hkl where $k \neq 3n$ ($n = 1, 2, 3, \dots$) suggesting loss of the crystallographic order and/or defects in the kaolinite structure caused by displacement type $\pm b/3$ or $\pm 2/3$ rotations. As the proportion of defects is reduced, it suggests the existence of $k \neq 3n$ with a profile similar to the hkl reflections (Aglietti and Porto Lopes 1986).

According to the Hinckley index, the structural disorder of kaolinite is more noticeable for Kao1 and Kao2 than for Kao3. The structural disorder suggests the existence of weathering products from feldspars, which are precursors of kaolinites well and/or poorly ordered mainly along the b axis. The Liètard index shows a high number of

TABLE I
Chemical composition of Kao1, Kao2 and Kao3 obtained by X-ray fluorescence.

Oxide	Kao1	Kao2	Kao3
	-----%-----		
SiO_2	66.4	53.3	52.2
Al_2O_3	24.7	35.0	37.0
$\text{SiO}_2/\text{Al}_2\text{O}_3$	2.70	1.52	1.41
Fe_2O_3	2.3	5.5	2.3
TiO_2	6.6	6.2	8.6

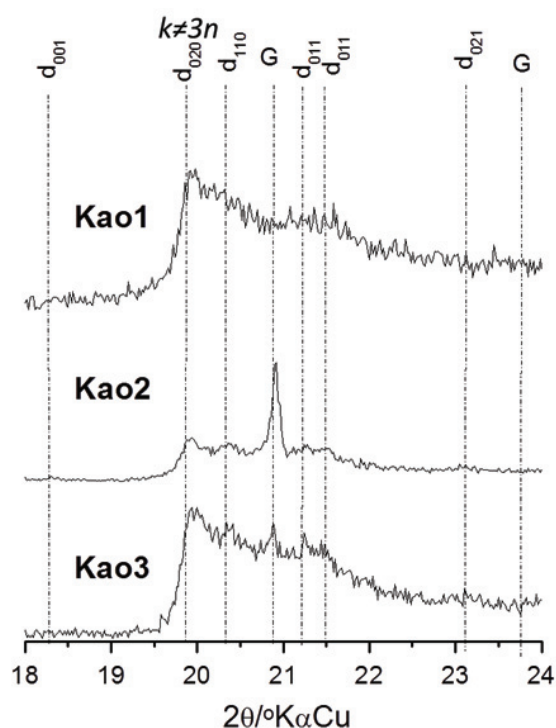


Figure 1 - X-ray diffraction patterns of kaolinites.

TABLE II
Crystallinity index of Kao1, and Kao2 Kao3.

Sample	Hinckley	Defect*	Liènard (R2)	Defect*
Kao1	0.549	Mean	0.026	High
Kao2	0.837	Mean	0.871	Mean
Kao3	0.465	High	0.453	High

*Categorized according to Aparicio et al. 2006.

defects for Kao1 and Kao3 and average for Kao2 obeying the following order: Kao1 > Kao3 > Kao2. Alterations in the kaolinite structure occur due to the isomorphic substitution by Fe^{3+} , spaces caused by the loss of Al^{3+} or electronic defects (Gaite et al. 1997, Aparicio et al. 2006). The chemical analysis indicating a high amount of Fe_2O_3 a loss of about 15.0 % Al_2O_3 and an increase of 20.0 % SiO_2 explain the presence of iron and spaces in the kaolinite structure of Kao1 as well as changes in the displacement $\pm b/3$ and rotation $\pm 2p/3$.

The FTIR spectra shows absorption bands of stretching vibrational OH groups (internal and external) present in the octahedral sites (Figure 2). At $3619\text{-}3621\text{ cm}^{-1}$ a band stretching vibrational absorption parallel to the *ab* plane is observed, i.e., in the common plane of apical oxygen of the tetrahedral sheet. The vibrational absorption bands in Kao1, Kao2, and Kao3 indicates a similarity to the internal OH groups of the three kaolinites. The presence of an absorption band at 3696 cm^{-1} characterizes perpendicular vibrational mode typical of the octahedral layer. It suggests a preservation of the OH groups from the kaolinite surface in all samples. It shows Kao1, Kao2, and Kao3 having their hydrogens linked with oxygen atoms bound to octahedral layer surface without considerable change. The vibrational mode of O-H bound to the atom of aluminum octahedral layer confirms what we had imagined.

However, the relationship of intensity of the stretch bands $3619\text{-}3621$ and 3696 cm^{-1} indicates different levels of pleochroism for Kao1, Kao2, and Kao3. Kao1 presents an inversion of the intensity of the bands, i.e., absence of pleochroism caused due to rotation $\pm 2\pi/3$, as shown in the XRD (Aglietti and Porto Lopez 1986). The ratio of intensity between the bands at 795 and 755 cm^{-1} also demonstrates loss of pleochroism. According to Deng et al. (2002), it occurs because of the particle size reduction. The absence of the 3652 cm^{-1} band in Kao3 shows vibrations directly related to weak hydrogen bonds typical of kaolinite of low crystallinity and the three vibrational modes remaining indicate OH groups coupled with moments of transition in the d_{001} plane (Farmer 2000, Frost et al. 2001). The shape of the band that corresponds to the angular-displacement vibration in 938 cm^{-1} indicates a kaolinite with a low number of hydrogen bonds in octahedral surface (Figure 2). The absorption band in 914 cm^{-1} relative to the angular-displacement vibration corresponds to the aluminol (Al-O-H) group. Furthermore, the O-H

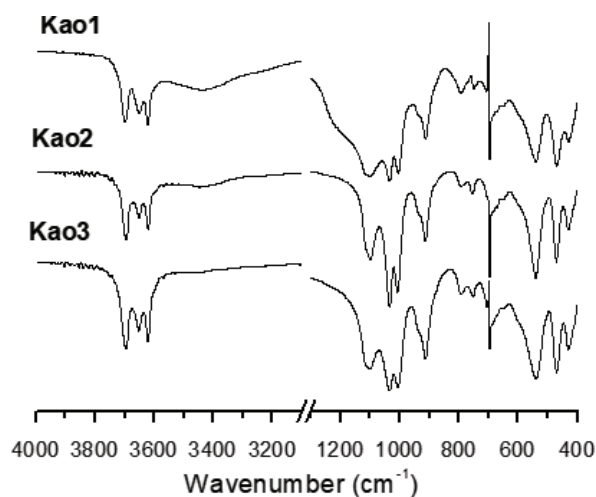


Figure 2 - FT-IR of the samples.

bond is unchanged for octahedral and tetrahedral sheets layers indicated by the stretching band at $3619\text{-}3621\text{ cm}^{-1}$.

In the 1010 cm^{-1} region, the Kao1, Kao2, and Kao3 are varied by absorption band of Si-O stretch suggesting different particle sizes as well as low crystallinity (Shoval et al. 1999). It is interesting to note that the concentrations of SiO_2 and Al_2O_3 are different for Kao2 and Kao3, which suggests two types of kaolinite. The relative intensities of 468 cm^{-1} (band angular vibration of Si-O) and 539 cm^{-1} (band angular vibration of Si-O-Si) showed that Kao1 has small amounts of Al_2O_3 suggesting the existence of isomorphic substitution in the octahedral site. Therefore, the changes in the vibrational modes detected by the loss of pleochroism, differences in the angular vibration suggest that kao1 has changed its crystalline structure due to isomorphic substitution.

The SEM of Kao1, Kao2 and Kao3 are typical of kaolinite (Figure 3). Kao1 presents groups and stacks in the form of pilasters and crystals in the form of lamellae. This structure called booklet with pictures and cylindrical vacancies, produces porosities in the kaolinite lattices while Kao2 presents rod-shaped forms of “stalactite”,

presenting better crystallinity. Kao3 shows strong placoid sheets, vertical growth toward the crystallographic c axis, showing mineral habit is not well defined. The mapping elements by EDS Kao1, Kao2 and Kao3 revealed that the predominant elements are Si, Al, O, Fe and Ti. The SEM shows that the morphology and particle sizes from Kao1, Kao2 Kao3 have several levels.

Mössbauer parameters and spectra at 298 K clearly show that Kao1, and Kao2 Kao3 possesses paramagnetic Fe^{3+} , suggesting isomorphic substitution of Fe^{3+} by Al^{3+} in the octahedral sites (Figure 4 and Table III). Furthermore, the widest line ($G = 0.62$ and 0.71 mm s^{-1}) indicate the presence of a triclinic kaolinite with layers influenced by differences in layer stacking of polytypes as well as by degree of layer stacking order. (Komusinski et al. 1981).

Although the Mössbauer spectra showed that the majority of kaolinites contain divalent iron in octahedral coordination, there were no values of $d = 1.130 \pm 0.050\text{ mm s}^{-1}$ for the Fe^{2+} substituted octahedral site. In turn, the orientation of the hydroxyl groups of octahedral sheets is determined by the stretching layer as well as by the symmetry of the crystal field in which the Fe^{3+} is located.

Mössbauer parameters found for Fe^{3+} are characteristic of isomorphic substitution in position *cis* and *trans-OH* (Taneja and Raj 1993). Therefore, it is not only the structural differences of kaolinites caused by defects and different octahedral shapes (Heller-Kallai and Rozenon 1981). Actually, the paramagnetic behavior observed reveals that kaolinite contains structural defects in its crystal lattice due to isomorphic substitution by Fe^{3+} , causing disorder and changing the spin-spin interaction making the paramagnetic (Gaite et al. 1997). Furthermore, the Mössbauer spectra are distinguishable showing polytypic modifications of kaolinite with different line-widths. The presence of Fe^{3+} forming OH-*cis* and OH-*trans* points out trapped hole centers in kaolinite structure,

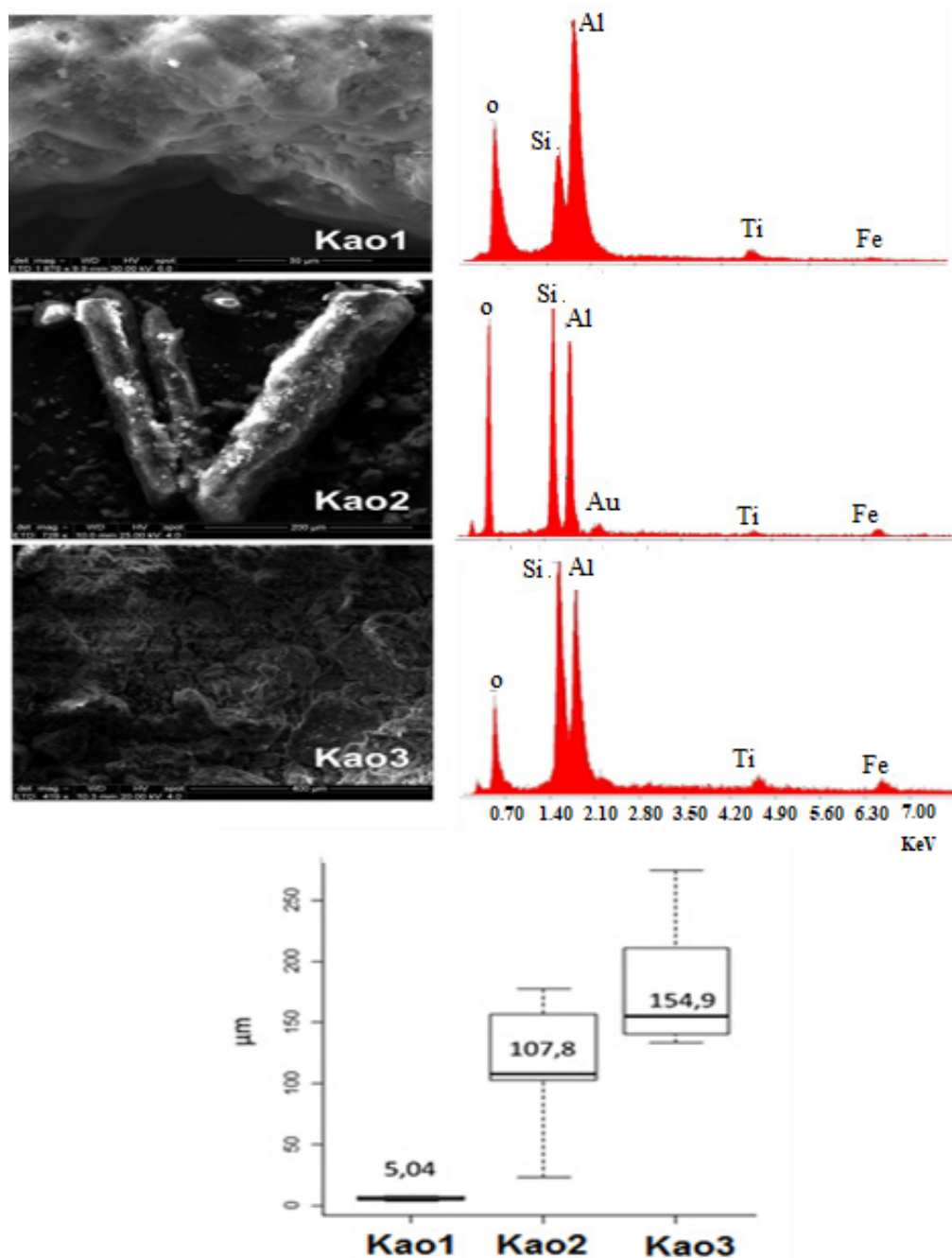


Figure 3 - Secondary electron images of scanning electron microscopy and EDS spectrum of the Kao1 and Kao2 Kao3 samples.

trapping locus for electron holes (holes), i.e., the substitution of Al^{3+} by Fe^{3+} forms a positive moiety of the trapped separated charge pair produced by electronic excitation (Coyne and Summers 1991). According to Sato et al. (2004), the holes in kaolinites occur in B- or C-sites which explains the formation of the chiral structure.

The relative areas of the fitted Mössbauer spectra indicate that isomorphic substitution is greater in the octahedral site in *cis* position, following this order of reason $\text{AR-}cis/\text{AR-}trans$: Kao1 (10.5) > Kao2 (4.0) > Kao3 (2.8). Mössbauer spectra *d* and *D* (*e*) of paramagnetic Fe^{3+} from silicate and/or iron oxides, not identified by previous physical techniques. The 80 K spectra demonstrate that only Kao2 shows low among iron oxides as well as a low supermagnetism effect (Rozenon et al. 1982). The Mössbauer parameters

(*d*, D/e e B_{hf}) are typical of hematite ($\alpha\text{-Fe}_2\text{O}_3$) and goethite ($\alpha\text{-FeOOH}$).

The study of the crystal structure observed for Kao1, Kao2 and Kao3 indicate a relation among the degree of substitution of Fe^{3+} at octahedral sites, characterized by the presence of Fe^{3+} at positions *cis/trans-OH* causing defects in the crystal structure of kaolinite, changing the crystalline arrangement and size of the average particle. Physical measurements reveal a larger amount of octahedrons consisting of Fe^{3+} at the position *cis* to Kao1, differing considerably in relation to Kao2 and Kao3.

The adsorption process of phosphorus is characterized by having the parameters of adjustments to the Freundlich isotherms values $R^2 \sim 0.99$ indicating that adsorption occurs in the presence of heterogeneous surfaces on octahedral sites (Table IV). The negative *n* values suggest low

TABLE III
Mössbauer parameters at 298 and 80K.

Samples	Site (^{57}Fe)	<i>d</i> (mm s^{-1})	<i>D</i> , <i>e</i> (mm s^{-1})	B_{hf} (T)	RA (%)
298 K					
Kao1	Fe^{3+} <i>cis</i>	0.355(5)	0.529(9)	0.0	91.3(2)
	Fe^{3+} <i>trans</i>	0.38(5)	1.6(1)	0.0	8.7(2)
Kao2	Fe^{3+} <i>cis</i>	0.351(2)	0.511(5)	0.0	80(2)
	Fe^{3+} <i>trans</i>	0.37(1)	0.93(4)	0.0	20(2)
Kao3	Fe^{3+} <i>cis</i>	0.355(3)	0.474(9)	0.0	69(2)
	Fe^{3+} <i>trans</i>	0.36(1)	0.93(3)	0.0	25(2)
	Fe^{3+}	0.37(3)	2.11(6)	-----	6.9(8)
80 K					
Kao1	Fe^{3+} <i>cis</i>	0.455(7)	0.59(1)	0.0	89(1)
	Fe^{3+} <i>trans</i>	0.69(3)	2.39(7)	0.0	11(1)
	$\text{Fe}^{3+}(\alpha\text{-Fe}_2\text{O}_3)$	0.48(2)	-0.20(2)	52.0(1)	8.0(8)
Kao2	$\text{Fe}^{3+}(\alpha\text{-FeOOH})$	0.45(1)	-0.23(2)	46.28(8)	30(1)
	Fe^{3+} <i>cis</i>	0.456(4)	0.550(6)	0.0	44.8(5)
	Fe^{3+} <i>trans</i>	0.56(1)	1.87(2)	0.0	16.9(6)
Kao3	Fe^{3+} <i>cis</i>	0.463(7)	0.59(1)	0.0	80(2)
	Fe^{3+} <i>trans</i>	0.60(3)	1.65(5)(3)	0.0	11(2)
	Fe^{3+}	0.47(4)	2.67(9)	0.0	9(2)

d = aFe isomer shift; *e* = isomer quadrupole splitting; *D* = Quadrupole; B_{hf} = hyperfine magnetic field; RA = relative area.

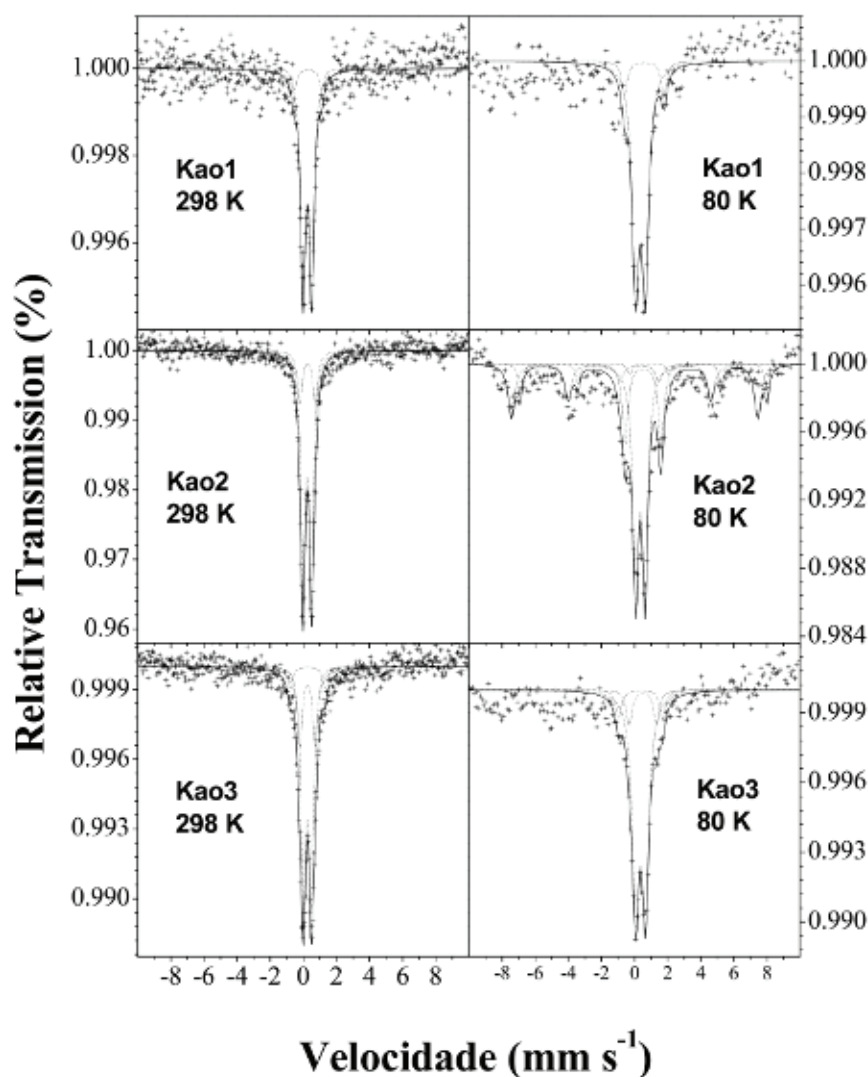


Figure 4 - Mössbauer spectra at 298K and 80K of the samples.

adsorption capacity of orthophosphate adsorption for all samples. It is explained by physical measurements mainly by substitution of Fe^{3+} , responsible for the differentiation of *cis*, and *trans* octahedral sites. The adsorption capacity obeys the following order $\text{Kao1} > \text{Kao2} > \text{Kao3}$. Largegren parameters (equation pseudo-second-order) clearly show that Kao1 retains faster orthophosphate on its surface. The increase in the substitution of Fe^{3+} -*cis* in the octahedral sites explains the retention of orthophosphate by Kao1 than Kao2 and Kao3. These findings reveal that the iron in the kaolinite

structure contributed with their chemical nature and high specific area as an important role in the transport and the removal of orthophosphate.

CONCLUSIONS

The mineral composition comprised kaolinite, quartz, gibbsite and goethite. As major element the EDS spectra and X-ray fluorescence showed the prevalence of Al and Si.

The presence of goethite indicates that Kao1 and Kao3 are not as weathering as Kao2. XRD

TABLE IV
Parameters of Freundlich adsorption kinetics and Largegren (equation pseudo-second-order).

Freundlich					
Amostra	K_f	n		R^2	$\text{Log } q_e = B + AX$
Kao1	17.333	-1.160		0.99	1.239 - 0.984 Log m
Kao2	17.954	-1.321		0.99	1.250 - 0.768 Log m
Kao3	18.192	-0.989		0.99	1.260 - 1.010 Log m
Largegren					
	q_c	k_2	H		t/q_t
Kao1	0.108	-2.534	-0.030	0.97	-33.803+ 9.256t
Kao2	0.174	-10.313	-0.310	0.99	-3.218+ 5.761
Kao3	0.183	-46.021	-1.550	0.99	-0.644+ 5.443

summarizes Kao1 and Kao2 as poorly crystallized/ordering, while kao2 as crystallized/ordering. As a result of average particle reduction, FTIR showed a higher inversion of the vibration bands relative to the hydroxyl groups in the range of 3619-3696 cm^{-1} and 795 and 755 cm^{-1} suggesting a natural loss of pleochroism of the Kao1 samples.

SEM Images showed that Kao2 is well-crystallized while Kao1 and Kao3, poorly crystallized. These results explain the values of the Hinckley, and Liétard indices obtained. Generally, samples substituted by Fe^{3+} form octahedral sites in the position *cis*-OH and *trans*-OH. The adsorption isotherms and kinetics, fitted by Freundlich and Largegren equations confirm the existence of heteronuclear sites on kaolinite surfaces

The Fe^{3+} -isomorphic substitution alters the magnetic properties of the kaolinite to paramagnetic according to the Mössbauer spectroscopy. The number of Fe^{3+} -OH-*cis* determines the capacity to adsorb orthophosphate.

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